SOLAR FLARE PARTICLE EFFECTS AND SEASONAL RADIOCARBON VARIATIONS IN TREE RINGS OF THE NORTHERN AND SOUTHERN HEMISPHERES

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ABSTRACT. Measurements of Δ^{14} C and δ^{13} C, separately in early wood and late wood of two Mediterranean pine species, with a nearly year-long growth season, were made to better define possible solar flare particle effects. The measurements were made for suitable periods around the years of big flares (*e.g.*, 1942, 1946, 1989) in *Pinus pinea* of the Adriatic area and *Pinus halepensis* from South Australia. We found that the ¹⁴C concentration is consistently greater in early wood than in late wood of the same growth ring and considerably higher in the Australian samples than in the European samples. The data so far reveal no evidence of solar-flare effect.

INTRODUCTION

Time-dependent variations in the ¹⁴C production rate are caused by changes in the galactic cosmicray flux as a consequence of solar modulation or by direct contribution from solar flares and nearby supernova explosions. As for solar-flare-induced variations, although high ¹⁴C values are possibly attributable to flare events before the nuclear testing period, they are obscured, for recent solar cycles, by the large increase in atmospheric ¹⁴C produced during the same period. At present, however, that contribution is greatly reduced, and the occurrence of large flare events comparable to that of February 23, 1956 (Shea 1990) encourages a search for their ¹⁴C effects (*cf.* Lingenfelter and Ramaty 1970; Castagnoli and Lal 1980). The study of ¹⁴C produced by solar flare particles, during periods when measurements of the energetic particle intensities and spectra, and observations of the Sun are available, plays a fundamental role in reconstructing past solar activity.

Following the investigations by Shea (1990) and Kostantinov *et al.* (1992), we present here a progress report on the search for ¹⁴C abundance variations in tree rings induced by solar-flare particles. We made and repeated the measurements of possible local enhancement of ¹⁴C production, both in the northern (NH) and southern hemispheres (SH), as an independent verification, because of the recent solar flares in September–October 1989, and checked for analogous effects from the large solar particle events of 1942 and 1946 (Pomerantz and Duggal 1974) (see Table 1).

METHODS

In the NH, we used wood samples for a pine tree (*Pinus pinea*) of the Ravenna pine forest located ca. 3 km from the Adriatic Sea, Italy (46°N); the samples were cut on 29 February 1992. In the SH, we took samples from a pine tree (*Pinus halepensis*) located at the Belair National Park near Adelaide, Australia (34°S); the samples were cut in September 1990. Branches and trunks were dissected to obtain disks 2–4 cm thick. Each disk was then cut into four sections. The parts were sanded

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on both faces in order to see a well-defined ring structure. We started from the outermost ring using a 10-mm woodworking chisel. Where the separation of early wood from late wood faded, we made a preliminary incision with a cutter, which made the chisel positioning easier. Wood chips were cut into match-like pieces; the excess wood was then rubbed away until the next ring of the sequence was reached.

Tree-Ring Samples		
Year	δ ¹³ C	$\Delta^{14}C$
(AD)	(‰)	(‰)
Belair		
90 LW*	-22.7	177.4
89 EW	-21.9	181.4
89 LW	-22.4	185.0
88 EW	-22.2	185.9
88 LW	-22.9	192.1
47 EW	-23.5	-11.7
47 LW	-23.5	-16.6
46 EW	-23.3	-19.1
46 LW	-21.1	-21.7
45 EW	-21.5	-20.1
45 LW	-21.5	-19.6
44 EW	-23.1	-35.1
Ravenna		
91 LW	-29.0	119.8
91 EW	-27.8	140.0
90 LW	-28.5	144.4
90 EW	-26.3	154.6
89 LW	-28.7	152.7
89 EW	-28.0	167.7
88 LW	-28.0	157.6

TABLE 1. Carbon Isotope Analyses in Early and Late Wood of Annual Tree-Ring Samples

*LW = late wood; EW = early wood

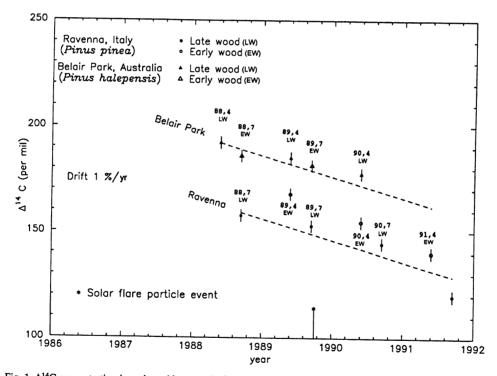
Chemical pretreatment consisted of boiling 50 g of wood in a 2M-KOH solution for 8 h. After a first waterflow washing, the sample was kept in a solution of 80 g NaClO₂ plus 110 ml of HCl-1M, in 3 liters of water, at 70–80 °C for 24 h. The sample was then drained, water washed, boiled in water for 3 h and rinsed until chlorides disappeared (0.1 N AgNO₃ test). The sample was then dried in an oven at 50–80 °C. This procedure isolates and purifies the cellulose fraction. *Ca.* 15 g of cellulose for each sample were oxidized in a quartz oven and the benzene was synthesized *via* Li₂C₂ and C₂H₂. Average efficiency of the cellulose to benzene transformation was *ca.* 93%. Teflon[®] vials of 7 ml, filled with 4.395 g of benzene and 75 mg of butyl-PBD, were used to determine ¹⁴C activity.

To minimize any possible "memory effect" of the line, we performed various syntheses of benzene from the samples: modern standard (ANU sucrose) samples, cellulose samples, followed by the modern standard again. Before their use, each vial was identified for background and efficiency. Beta counting was done by a Wallac Quantulus 1220^{TM} ; samples and standards were measured *quasi-simultaneously* for *ca*. 5 days for each vial. During counting, the standard quench parameter (external source) was continuously monitored; the endpoint spectrum variation was also monitored. It remained within three channels of the ¹⁴C endpoint. These small endpoint variations were corrected every time by adjusting the counting window appropriately.

Because we checked seasonal Δ^{14} C variations, we considered separately the statistical and systematic components of the uncertainty. We evaluated the reproducibility of our measurements using four modern standards, which were oxidized during the sample analyses. The four activity values obtained for the modern standard, each having a 2‰ statistical uncertainty, show that the measurements became statistically self-consistent if uncertainty was multiplied by the factor 1.4. Because the statistical counting uncertainty for each sample was not >2‰, the *random* error became 2.8‰. The systematic component of the uncertainty was 1.5‰, and the overall uncertainty was 3.2‰.

RESULTS AND CONCLUSION

Figure 1 shows Δ^{14} C results for the Adelaide and Ravenna measurements in 1988–1991. For both sites, we made the measurements in early wood and late wood separately, as late wood-tracheids are supposedly produced in the late season (Attolini *et al.* 1990), *i.e.*, after September-October-November in the Ravenna region. As can be seen in both graphs, the overall annual decreasing trend is consistent with the decrease of 10‰ per year observed in the atmosphere in various stations of the NH, as reported by Levin *et al.* (1989). This is mainly due to nuclear test effects and to the Suess effect from fossil-fuel combustion.



The Δ^{14} C values of the Australian samples, *ca.* 2% higher than the Ravenna samples, appear to be real, as the measurements were made simultaneously. On the contrary, Barbetti *et al.* (1992) for AD 1600–1800, Lerman, Mook and Vogel (1970) and Vogel, Fuls and Visser for AD 1835–1900 reported a Δ^{14} C greater in the NH than in the SH. On the other hand, our Δ^{14} C measurements (see Fig. 2)

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before the bomb-test period agree with the values found by Stuiver and Quay (1981) and Vogel, Fuls and Visser (1993). Suess-effect measurements in 1950 by the same authors seem to indicate a smaller fossil dilution in the SH (15‰) than in the NH (20‰). However, it is difficult to attribute the difference we found only to this cause. The present time effect might be due to ¹⁴C from nuclear explosions, occurring at different times in the SH and NH (see Tans, de Jong and Mook 1979; Olsson and Possnert 1992; Dai *et al.* 1992), and to its removal time constant, of *ca.* 25 yr (see Dai, Qian and Fan 1992). To verify this possibility, we plan to make more measurements using ancient dated wood, and compare them with the southern atmospheric ¹⁴C concentration.

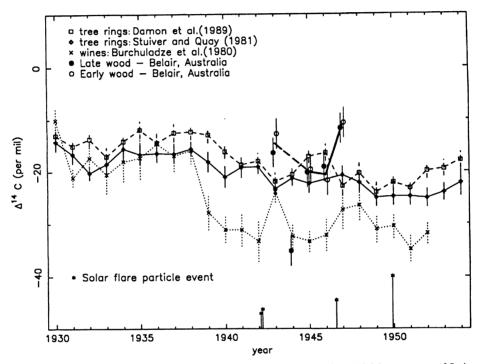


Fig. 2. Pre-bomb Δ^{14} C concentration in early and late wood of tree rings of the SH. Measurements of Stuiver and Quay (1981) and Damon, Cheng and Linick (1989) in annual tree rings and of Burchuladze *et al.* (1980) in wines are shown for comparison. The segment height of major solar flare particle events is proportional to the total flux enhancement (Pomerantz and Duggal 1974).

For the Ravenna pine, the late-wood values of ¹⁴C appear to be systematically lower than adjacent early wood. We also plan more investigations to verify if this effect could be caused by the area's proximity to a large, populated area, as suggested by the seasonal variation of CO_2 concentration reported in Levin *et al.* (1989). Whether this effect is real or not, the coupling of early with late wood values provides an upper limit of the uncertainty of reproducibility.

In Figure 2, we also plotted the annual measurements of Stuiver and Quay (1981), Damon, Cheng and Linick (1989) and those from wines of Burchuladze *et al.* (1980). Our results seem a bit higher, but we cannot speak definitively of systematic differences. Although Figures 1 and 2 show significant variations between two subsequent measurements, none can be attributed conclusively to the solar-flare effect. However, for the solar flare years, 1942–1946, there seems to be an enhancement of the expected order of magnitude, according to Lingenfelter and Ramaty (1970). The time interval covered by our measurements is too small to allow an evaluation of solar modulation effects when

comparing our results with sunspot index Rz or the magnetic index AA, as shown in Figure 3. We plan to continue the investigation, making more Δ^{14} C measurements to find possible effects related to flares before 1950. We also plan to clearly separate variations induced by solar activity and the Suess effect that may obscure the changes in ¹⁴C production due to solar flares.

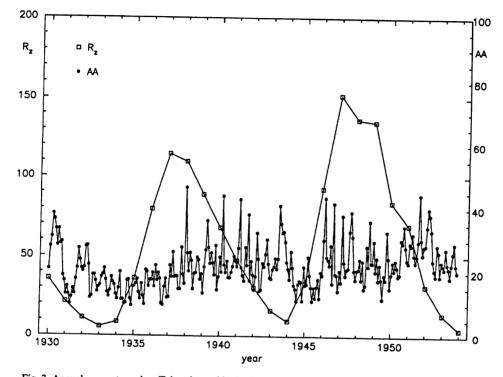


Fig. 3. Annual sunspot number (Rz) and monthly AA index for the period AD 1930-1954 (Mayaud 1973)

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